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Hällristning
Fiskare från
bronsåldern

Rock carving
Bronze age
fishermen



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Hydrografiska avdelningen, Göteborg

Tidal and Spectral Analyses of Kattegat
Time Series of Current and Salinity

by

Henryk Bieler and Artur Svansson

Januari 1978

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1. Introduction.

As a part of a project to determine the transport of water and matter in the northern Kattegat, automatically recording current meters, type Aanderaa, were anchored at various times and places. This paper deals with data recorded during the period 1975-04-19 -- 1975-08-19. The measuring depths were 15 m and 30 m.

Fig. 1 shows the area with depth topography. The position of the current meters was N 57°32' E 11°19.5' near the position of the former Läso Nord lightship.

2. Observation Material.

Current direction and velocity were recorded every 10th minute on magnetic tape. After a first processing the data were checked and corrected. In a very few cases empty spaces had to be filled with linearly interpolated values. In a second processing N- and E-components were computed and stored on a magnetic disc.

Hourly mean values centered at whole hours were computed of both N- and E-components respectively. This procedure may introduce noise; a more sophisticated filter technique had probably been a better choice. The data sets used for the tidal and spectrum analysis consisted of these hourly values.

Two periods were investigated, a shorter one (I), 75-04-19 -- 75-06-16 and a longer one (II), 75-04-19 -- 75-08-19, the latter for the 30 m depth only however. The mean currents were as follows:

	N-comp	E-comp
15 m I	-0.06 cm/s	-8.62 cm/s
30 m I	-3.77 cm/s	-1.65 cm/s
30 m II	-2.90 cm/s	-2.35 cm/s

3. Tidal analysis.

3.1. Method.

A copy of the analysis program TIFA (Murray 1964) was kindly placed at our disposal by Dr. G. Lennon, IOS, Bidston.

The program computes the amplitude H and GMT-phase G for Doodson's (1928) constituents. For period I, 26 constituents, for period II, 28 constituents were determined.

3.2. Results.

Tables 1-3 indicate the constituents which were used in TIFA. Whereas computation I covering 1415 hours is somewhat short for the long periods Mf and Msf, these two constituents could be included in II. A comparison between the remaining constituents computed both in I and II shows small differences. Note the phase coincidence in P1 and K1 and also in the T2, S2, K2 group. Due to the shortness of the data time series (to separate T2 and S2 for instance a complete year is required) and in the absence of other additional information, relationships of the equilibrium tides for amplitudes and identical values for phase was used.

Figs 2 and 3 show the M2 tidal ellipses for 15 m and 30 m respectively. Surprisingly they have opposite direction of rotation, 15 m cum sole and 30 m contra solem.

The tidal currents in the Kattegat have been analyzed before. Jacobsen (1913) presented harmonic constants for M2 at L/V:s Anholt Knob (3 month's data series) and Schultz's Grund (one year) of currents measured 6 times a day at 7 depths by means of a pendulum device. Jacobsen only analyzed the longitudinal component. His determination of eddy viscosity coefficients by means of the vertical distribution of the tidal currents is a classical work. A similar time series measured at L/V Läsö Rende covering the period 1912-09-05 -- 1913-11-14 was not

processed until 1968 then by Rossiter. As L/V Lässö Rende was situated in a narrow passage the currents are highly longitudinal, for M2 as follows:

depth m	H cm/s	G°
2.5	20.4	321
5	22.6	321
10	26.0	330
15	25.9	326
20	23.2	318

The vertical variations are small in contrast to conditions at L/V Anholt Knob where the surface (2.5 m) current is twice the value at 25 m depth. Defant (1934) compares measurement with results of a barotropic canal model (M2). The computed phase is approx. 325° at the border between the Skagerrak and the Kattegat and increases southwards (330° at the position of our current meter and 325° at the position Lässö Rende).

4. Spectral Analysis.

By means of power spectrum analysis the energy distribution as a function of frequency (hours^{-1}) was computed for the current time series used above. The tidal components computed by TIFA were subtracted before the analysis. The spectrum was determined according to Blackman and Tukey (1958), with small modifications as described in Mälkki (1975).

A program SPC computes autocovariance and spectrum of the time series at issue. The power spectrum $P(f)$ is a Fourier transform of the autocovariance function $C(\tau)$. For a continuous case we have:

$$P(f) = \int_{-\infty}^{\infty} C(\tau) \cdot e^{-i2\pi f\tau} d\tau \quad -\infty \leq f \leq \infty$$

$$C(\tau) = \int_{-T}^{T} P(f) \cdot e^{i2\pi f\tau} df \quad -T \leq \tau \leq T$$

where $f = \text{frequency } (\text{hour}^{-1})$

$\tau = \text{time lag } (\text{hour})$

For a discrete case is:

$$P(f) = \Delta t \sum_{\tau=1}^{N-1} c(\tau) \cdot e^{-i2\pi f \Delta \tau}$$

4.1. Computer Program.

Estimation is made of $P(f)$ associated with a time series computed by forming autocovariances for various time lags and obtaining their Fourier transforms. In order to smooth out $P(f)$ and $C(\tau)$ the following estimations were used:

$$C(\tau) = 1/N \sum_{t=1}^{N-1} (x_t - \bar{x})(x_{t+\tau} - \bar{x})$$

$$P(f) = 2 \cdot \Delta t \cdot [C(0) + \sum_{\tau=1}^{m-1} v(\tau) C(\tau) \cdot e^{-i2\pi\tau/m}]$$

where

$N = \text{total number of observations used}$

$$\bar{x} = 1/N \cdot \left(\sum_{j=1}^N x_j \right) = \text{mean value of all observations}$$

$$\tau = j \Delta t \quad 0 \leq \tau \leq m \Delta t$$

$v(\tau) = \text{window}$

$f = j/2m \Delta t = \text{frequency interval}$

$t = \text{time step between two observations}$

$m = \text{number of lags}$

4.2. Data Set.

The data set consists of a time series. The following information is needed:

N , Δt and m . Furthermore a window $v(\tilde{\tau})$ is chosen as a Tukey (Hanning) window:

$$v(\tilde{\tau}) = 0.5(1 + \cos(\pi\tilde{\tau}/m))$$

Blackman and Tukey (1958) show that the application of this window widens the main lobe of the equivalent analytical filter and reduces the side. Random errors in the data will introduce uncertainties into the estimates of $P(f)$ which may be determined in terms of 95 % confidence limits. These limits have been obtained to depend upon the v (number of degrees of freedom) defined as:

$$v = 2N/m b \quad \text{where } b \text{ is bandwidth of Tukey window.}$$

4.2. Method.

Spectra can be very effectively computed by a "cyclic" method according to Cartwright (Murray 1964):

$$w(\tilde{\tau}) = v(\tilde{\tau}) \cdot c(\tilde{\tau}) + 2w(\tilde{\tau} + 1) \cdot \cos(\pi\tilde{\tau} j/m) - w(\tilde{\tau} + 2)$$

$$w(m) = 0$$

$$w(m-1) = c(m-1) - v(m-1)$$

If we use the repetitive cycle ($m-1$) times, it finally yields $w(1)$ and $w(2)$, which can be shown to satisfy the relation

$$w(1) - w(2) \cdot e^{-i\pi\tilde{\tau}/m} = v(\tilde{\tau}) \cdot c(\tilde{\tau}) \cdot e^{-i\pi\tilde{\tau}/m}$$

4.3. Results.

Fig. 4 presents an example of the current data, on top in original and on bottom with TIFA computed tides subtracted, the last one used for spectrum analysis.

Fig. 5 shows the result of an analysis of the long series II East Comp at 30 m depth. There are certain peaks above the confidence limits at the following middle frequency periods: 12.1, 14.2, 17.2 and 21.7 hours.

The peak of 14.2 hours clearly originates from an inertial current. The high frequency range has a characteristic general slope. This slope, here -2 is often compared with Ozmidow's -5/3 slope in the classical theory of 3-dimensional turbulence:

$$E(k) = a \epsilon^{2/3} k^{-5/3}$$

where

a = constant

ϵ = rate of energy dissipation

k = scalar wave number

Fig. 6 shows the result of the power spectrum analysis of period I, N-component of 15 m and 30 m and Fig. 7, period I, E-component of 15 m and 30 m. Note that the point density is smaller than in Fig. 5.

Also salinities measured simultaneously were processed with the power spectrum technique, Fig. 8. As no TIFA-pretreatment was carried out with this data set, a comparison with correspondingly untreated current data was assumed to be more consistent, see Fig. 9.

Acknowledgement

We express thanks to Dr. G. Lennon for placing TIFA at our disposal and also for his valuable comments on the computational results.

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Fig. 1

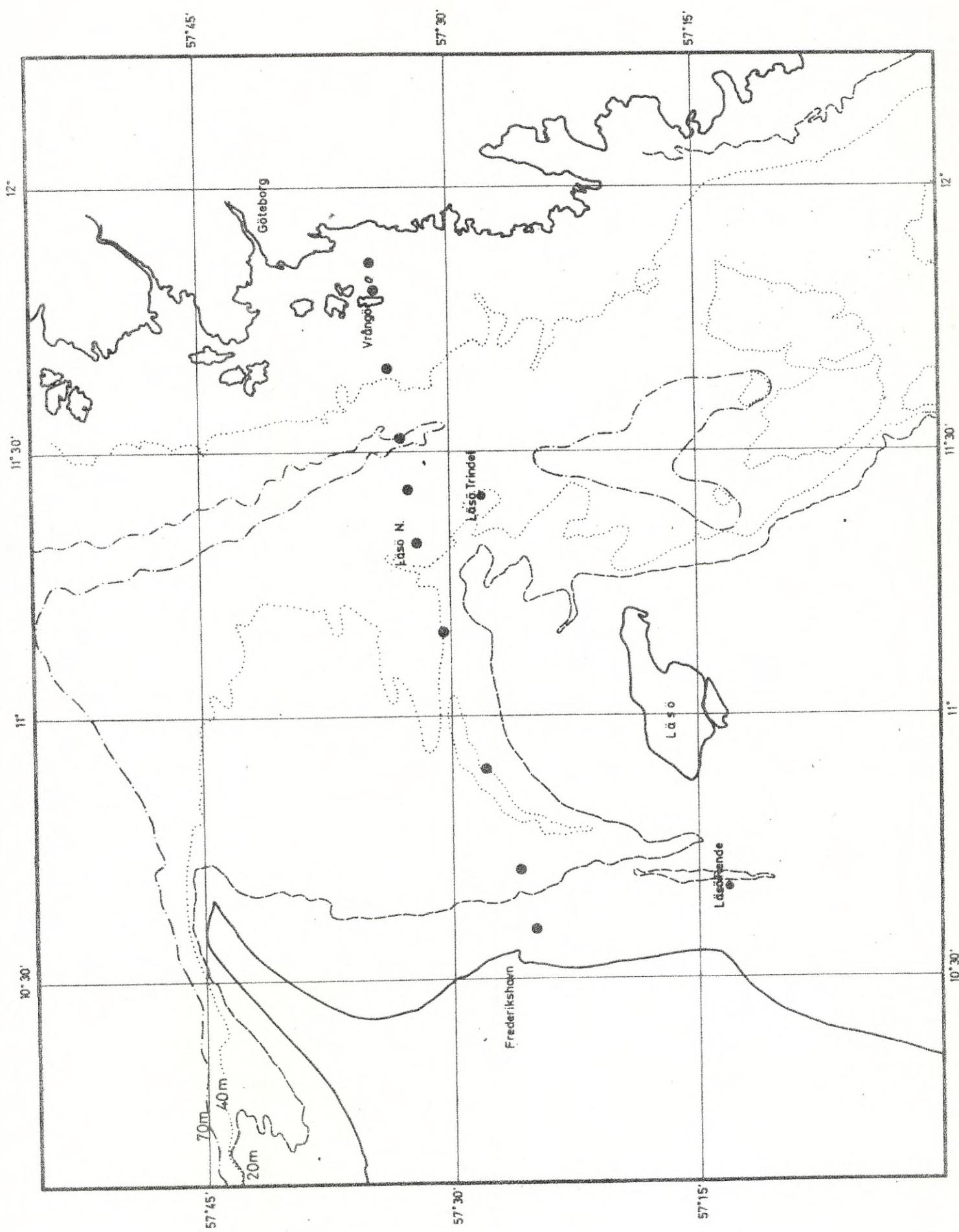
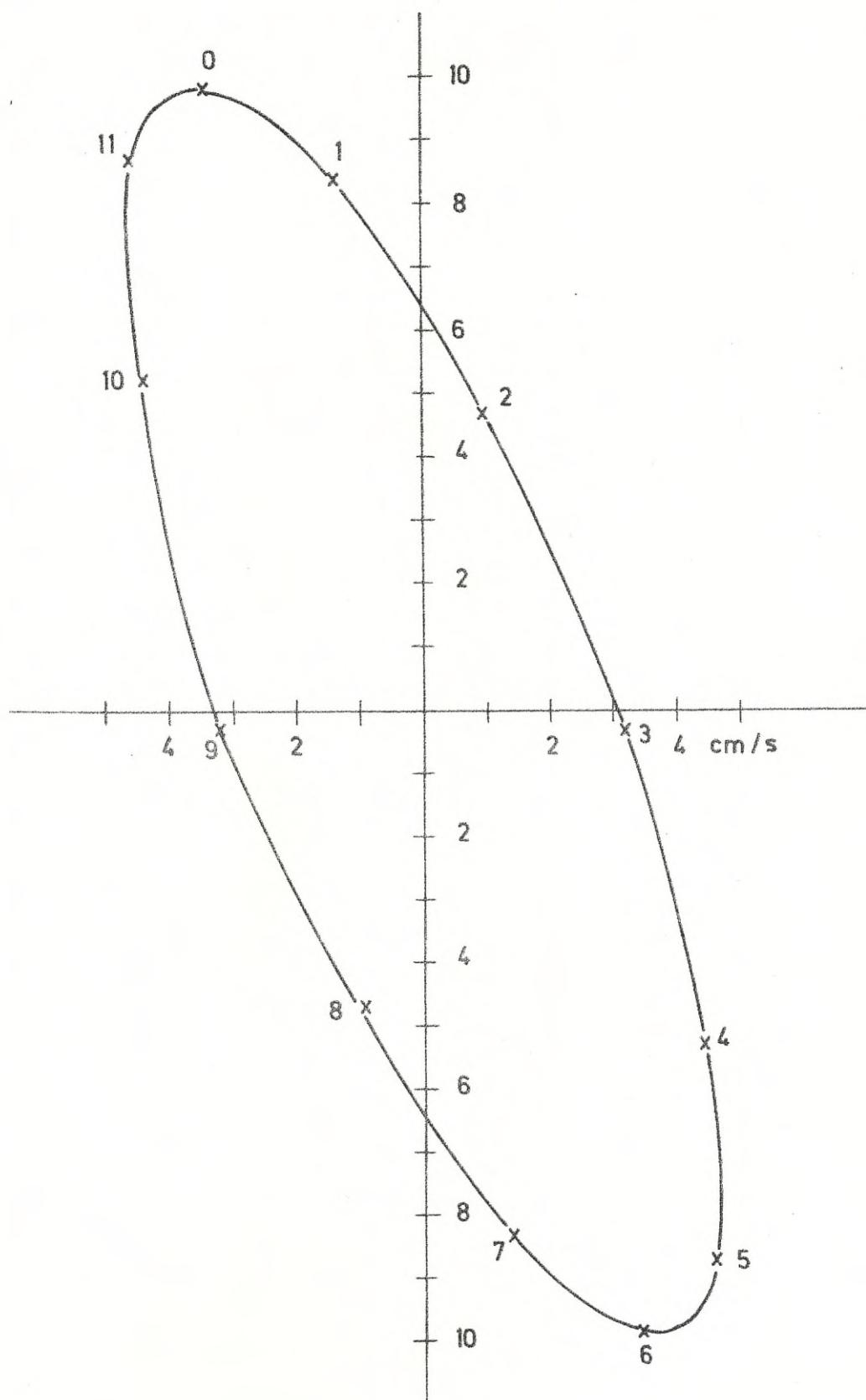
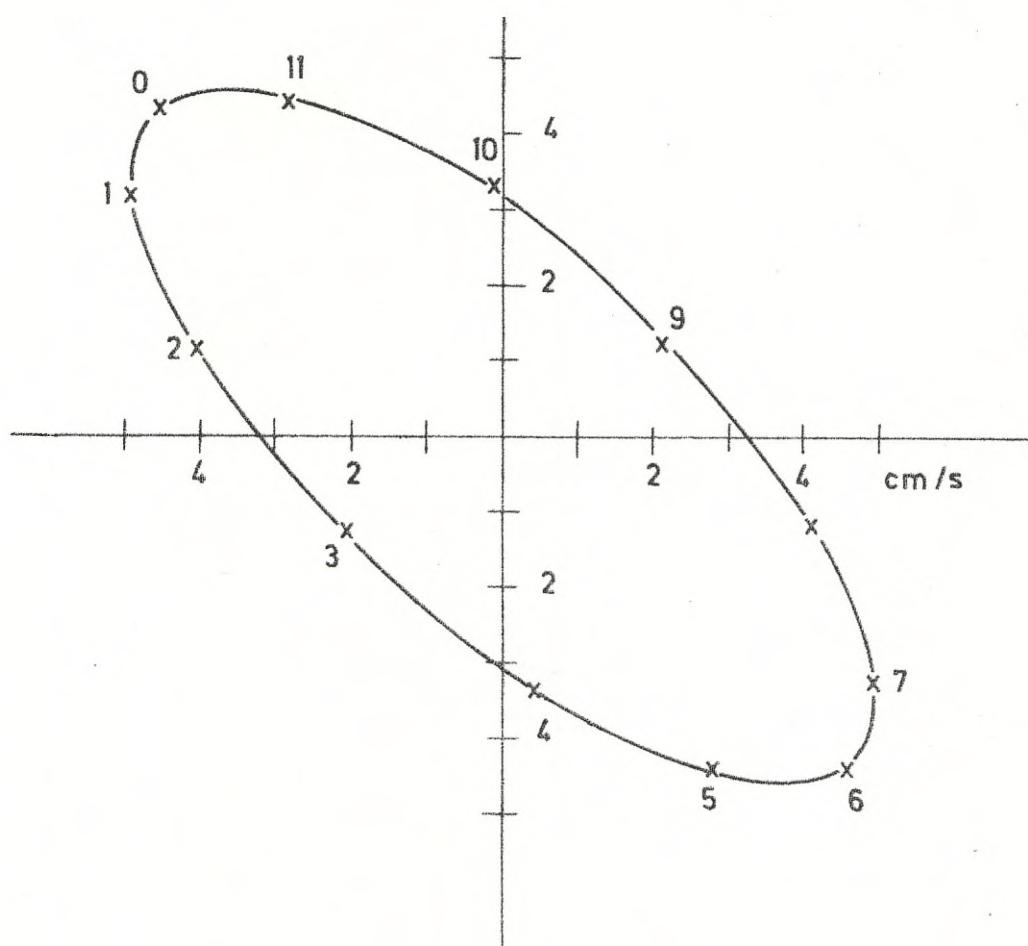


Fig. 2



M 2 Tidal Ellipse 15 m Computational period 75 04 19 - 75 06 16

Fig. 3



M 2 Tidal Ellipse 30 m Computational period 75 04 19 - 75 06 16

Fig. 4

Hourly Means of Current Meter Data, N 57° 32' E 11° 10.5', 15 m Depth

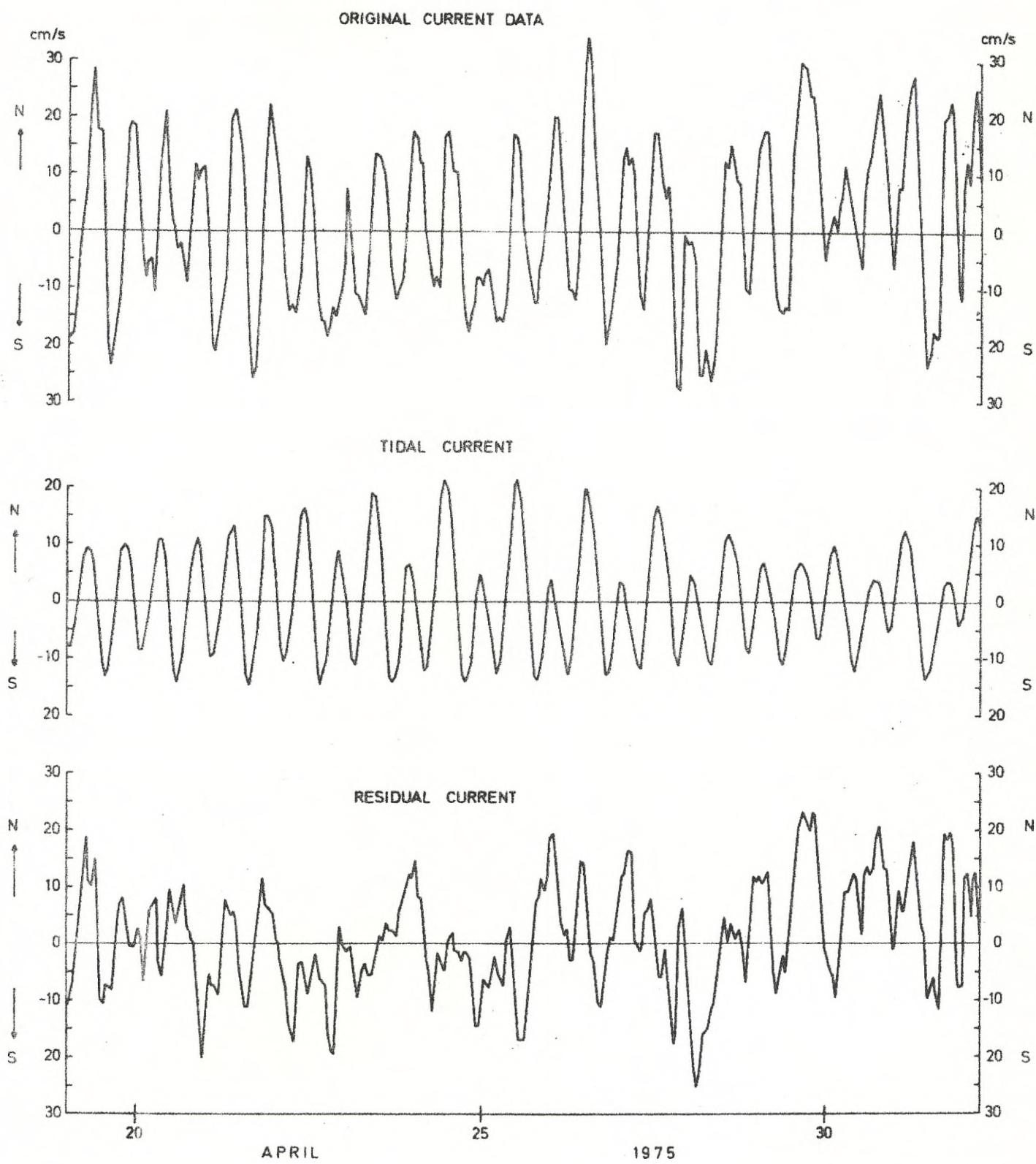
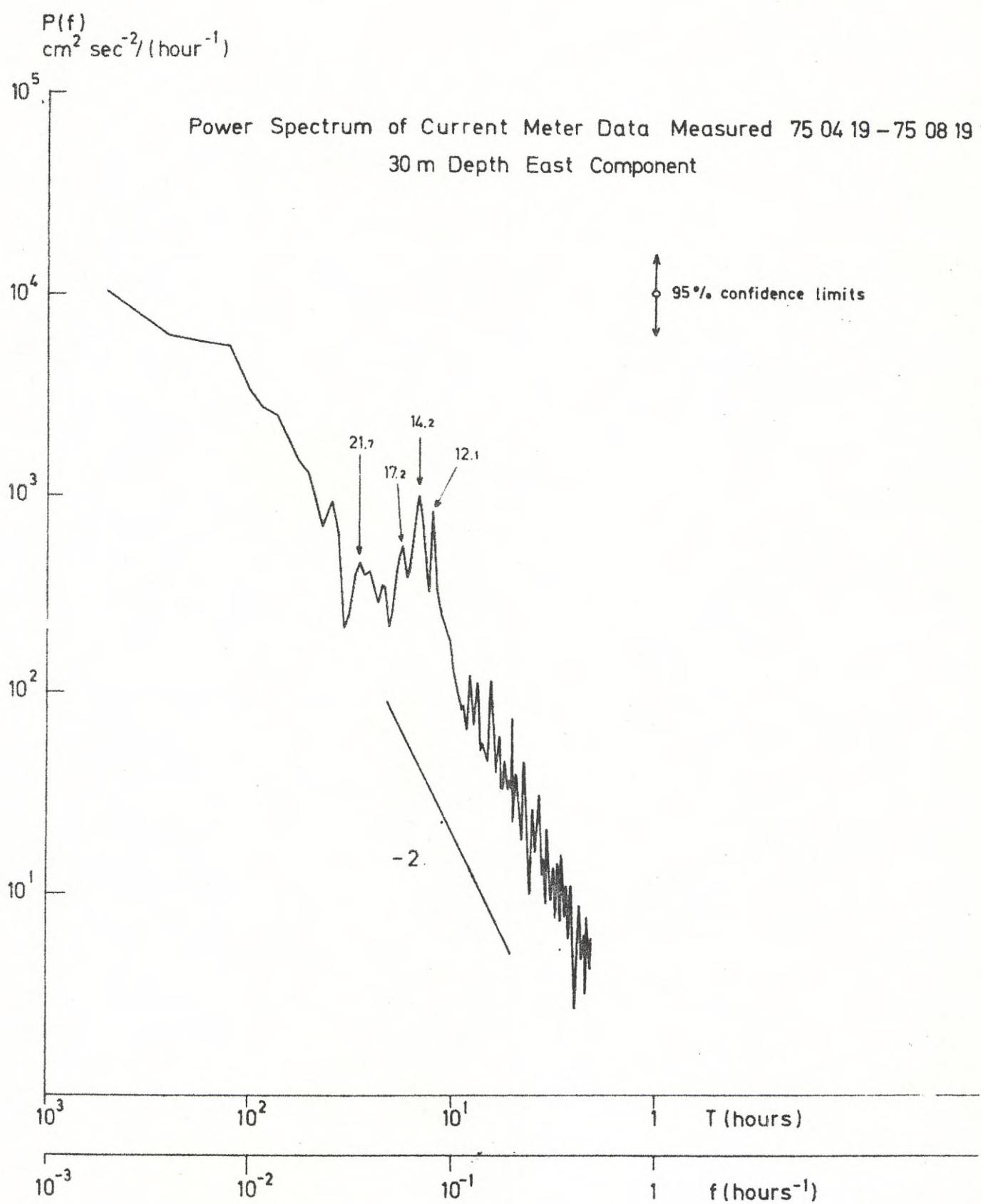


Fig. 5



$P(f)$
 $\text{cm}^2 \text{s}^{-2}/(\text{h}^{-1})$

Power Spectrum of Current Meter Data Measured
75 04 19 - 75 06 16

North Component

— 15 m
- - - 30 m

↑ 95% confidence limits

10^2

10^3

10^1

10^2

10^1

T_f
6

$f (\text{hours}^{-1})$

10^0

10^{-1}

10^{-2}

10^{-3}

Fig. 7

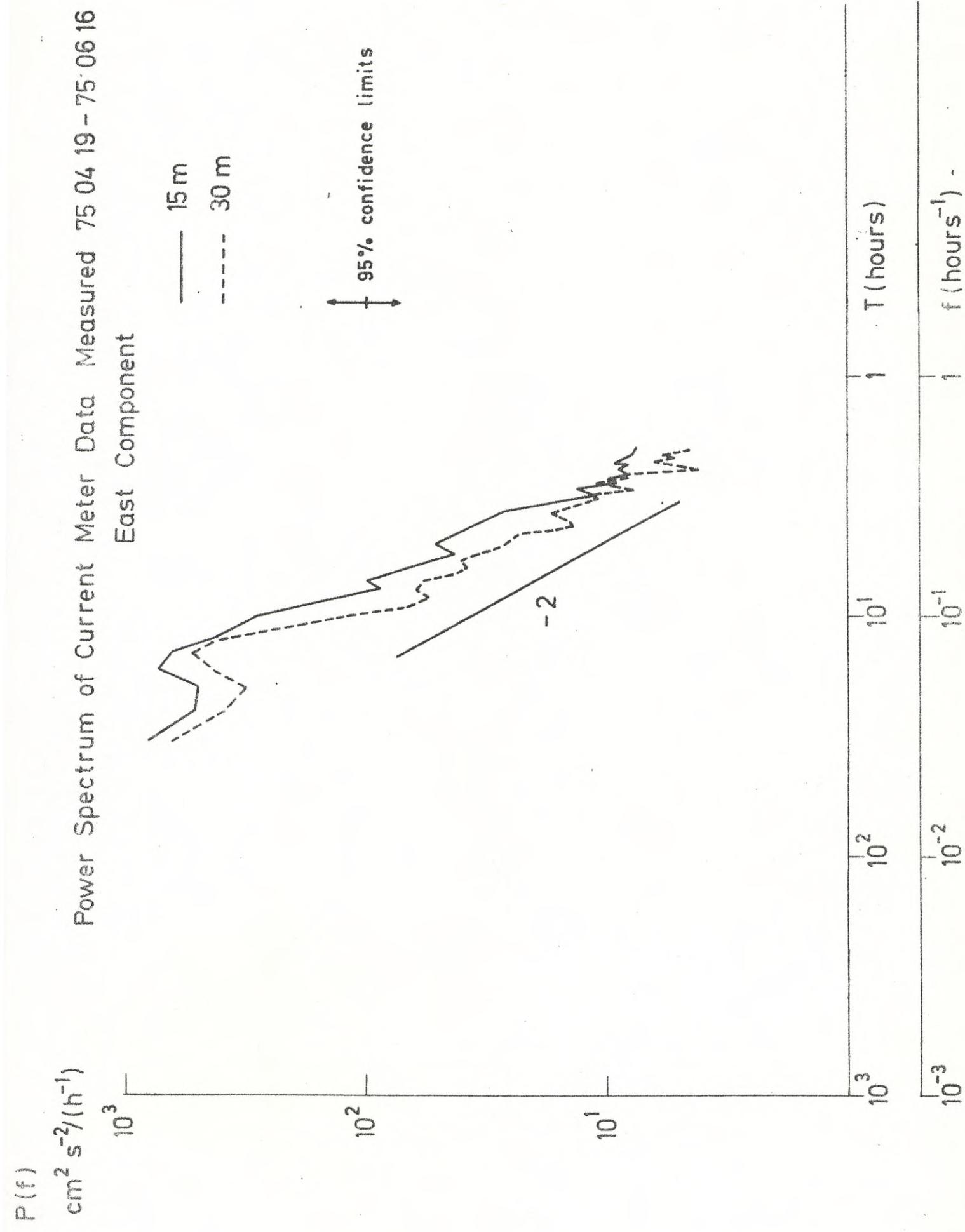


Fig. 8

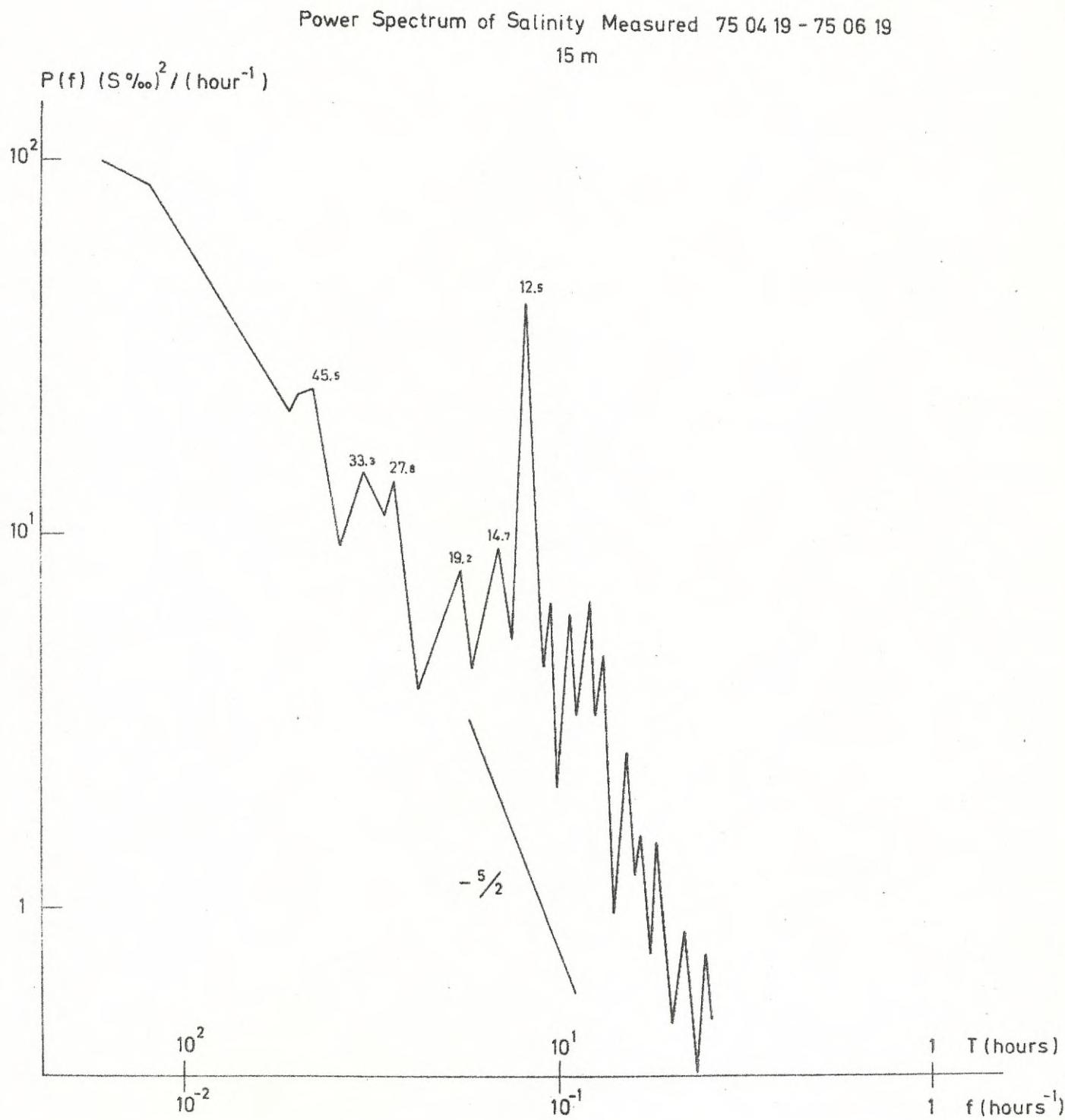


Fig. 9

Power Spectrum of Current Meter Data Measured 75 04 19 - 75 06 19
Tides included
N - Component 15 m

